

UC Irvine

UC Irvine Previously Published Works

Title

Three decades of global methane sources and sinks

Permalink

<https://escholarship.org/uc/item/04p4q8dd>

Journal

Nature Geoscience, 6(10)

ISSN

1752-0894

Authors

Kirschke, Stefanie
Bousquet, Philippe
Ciais, Philippe
et al.

Publication Date

2013-10-01

DOI

10.1038/ngeo1955

Copyright Information

This work is made available under the terms of a Creative Commons Attribution License, available at <https://creativecommons.org/licenses/by/4.0/>

Peer reviewed

Three decades of global methane sources and sinks

Stefanie Kirschke *et al.**

Methane is an important greenhouse gas, responsible for about 20% of the warming induced by long-lived greenhouse gases since pre-industrial times. By reacting with hydroxyl radicals, methane reduces the oxidizing capacity of the atmosphere and generates ozone in the troposphere. Although most sources and sinks of methane have been identified, their relative contributions to atmospheric methane levels are highly uncertain. As such, the factors responsible for the observed stabilization of atmospheric methane levels in the early 2000s, and the renewed rise after 2006, remain unclear. Here, we construct decadal budgets for methane sources and sinks between 1980 and 2010, using a combination of atmospheric measurements and results from chemical transport models, ecosystem models, climate chemistry models and inventories of anthropogenic emissions. The resultant budgets suggest that data-driven approaches and ecosystem models overestimate total natural emissions. We build three contrasting emission scenarios — which differ in fossil fuel and microbial emissions — to explain the decadal variability in atmospheric methane levels detected, here and in previous studies, since 1985. Although uncertainties in emission trends do not allow definitive conclusions to be drawn, we show that the observed stabilization of methane levels between 1999 and 2006 can potentially be explained by decreasing-to-stable fossil fuel emissions, combined with stable-to-increasing microbial emissions. We show that a rise in natural wetland emissions and fossil fuel emissions probably accounts for the renewed increase in global methane levels after 2006, although the relative contribution of these two sources remains uncertain.

Reconstructions of atmospheric methane (CH_4) concentrations between 1850 and the 1970s have been made using air trapped in polar ice cores and compacted snow. The data reveal an exponential increase in CH_4 levels in the atmosphere from 830 ppb to 1500 ppb in the late 1970s¹. Direct measurements of CH_4 in the atmosphere began in 1978², and reached global coverage after 1983. Today, CH_4 concentrations can be assessed using discrete air samples collected regularly at the surface, continuous measurements made at the surface^{2–6} or in the troposphere^{7–9}, and remotely sensed measurements of atmospheric CH_4 columns retrieved from the surface or from space^{10–12} (see Supplementary Section ST1). Surface-based observations from four networks (National Oceanic and Atmospheric Administration, NOAA¹³; Advanced Global Atmospheric Gases Experiment, AGAGE¹⁴; Commonwealth Scientific and Industrial Research Organization, CSIRO⁵; and University of California Irvine, UCI¹⁵) show consistent changes in the global growth rate of annual CH_4 concentrations since 1980 (Fig. 1 and Supplementary Section ST1). The agreement between these networks has improved with increasing coverage. The standard deviation for the global annual growth rate decreased from ± 3.3 ppb yr^{-1} in the 1980s to ± 1.3 ppb yr^{-1} in the 2000s. These data reveal a sustained increase in atmospheric CH_4 levels in the 1980s (by an average of 12 ± 6 ppb yr^{-1}), a slowdown in growth in the 1990s (6 ± 8 ppb yr^{-1}), and a general stabilisation from 1999 to 2006 to 1773 ± 3 ppb. Since 2007, CH_4 levels have been rising again¹⁴, and reached 1799 ± 2 ppb in 2010. This increase reflects a recent imbalance between CH_4 sources and sinks that is not yet fully understood¹³.

Previous reviews of the global CH_4 budget have focused on results from a few studies only^{13,16–19}. These studies covered different time windows and employed different assumptions, making it difficult to interpret the decadal changes presented. Only very few studies addressed multi-decadal changes in CH_4 levels^{20,21}. Here we construct a global CH_4 budget for the past three decades by combining bottom-up and top-down estimates of CH_4 sources and the chemical CH_4 sink (Box 1). We use chemical transport models — constrained by atmospheric CH_4 measurements — to estimate CH_4 fluxes using top-down atmospheric inversions. We compare these

fluxes with those simulated by ecosystem models of wetland and biomass burning emissions and by data-driven approaches for other natural sources (Methods and Supplementary Section II). We also gather recent data from fossil fuel CH_4 emission inventories based on energy use statistics, and from agricultural and waste inventories based on livestock and rice paddy statistical data.

Sources and sinks

The global atmospheric CH_4 budget is determined by many terrestrial and aquatic surface sources, balanced primarily by one sink in the atmosphere. CH_4 emissions can be broadly grouped into three categories: biogenic, thermogenic and pyrogenic. Biogenic sources contain CH_4 -generating microbes (methanogens)¹⁷, and comprise anaerobic environments such as natural wetlands and rice paddies, oxygen-poor freshwater reservoirs (such as dams), digestive systems of ruminants and termites, and organic waste deposits (such as manure, sewage and landfills). Thermogenic CH_4 , formed over millions of years through geological processes, is a fossil fuel. It is vented from the subsurface into the atmosphere through natural features (such as terrestrial seeps, marine seeps and mud volcanoes), and through the exploitation of fossil fuels, that is, through the exploitation of coal, oil and natural gas. Pyrogenic CH_4 is produced by the incomplete combustion of biomass and soil carbon during wildfires, and of biofuels and fossil fuels. These three types of emissions have different isotopic $\delta^{13}\text{C}$ signatures ($\delta^{13}\text{C} = [({}^{13}\text{C}/{}^{12}\text{C})_{\text{sample}}/({}^{13}\text{C}/{}^{12}\text{C})_{\text{standard}}] - 1) \times 1000$): -55 to -70‰ for biogenic emissions, -25 to -55‰ for thermogenic emissions, and -13 to -25‰ for pyrogenic emissions^{20,22,23}. The isotopic composition of atmospheric CH_4 — measured at a subset of surface stations — has therefore been used to constrain its source^{20–24}. CH_4 emissions by living plants under aerobic conditions do not seem to play a significant role in the global CH_4 budget (Supplementary Section ST8); some very large²⁵ estimates of this source published in 2006 have not been confirmed²⁶.

The primary sink for atmospheric CH_4 is oxidation by hydroxyl radicals (OH), mostly in the troposphere, which accounts for around 90% of the global CH_4 sink. Additional oxidation sinks include methanotrophic bacteria in aerated soils^{27,28} ($\sim 4\%$), reactions with

*A full list of authors and their affiliations appears at the end of the paper.

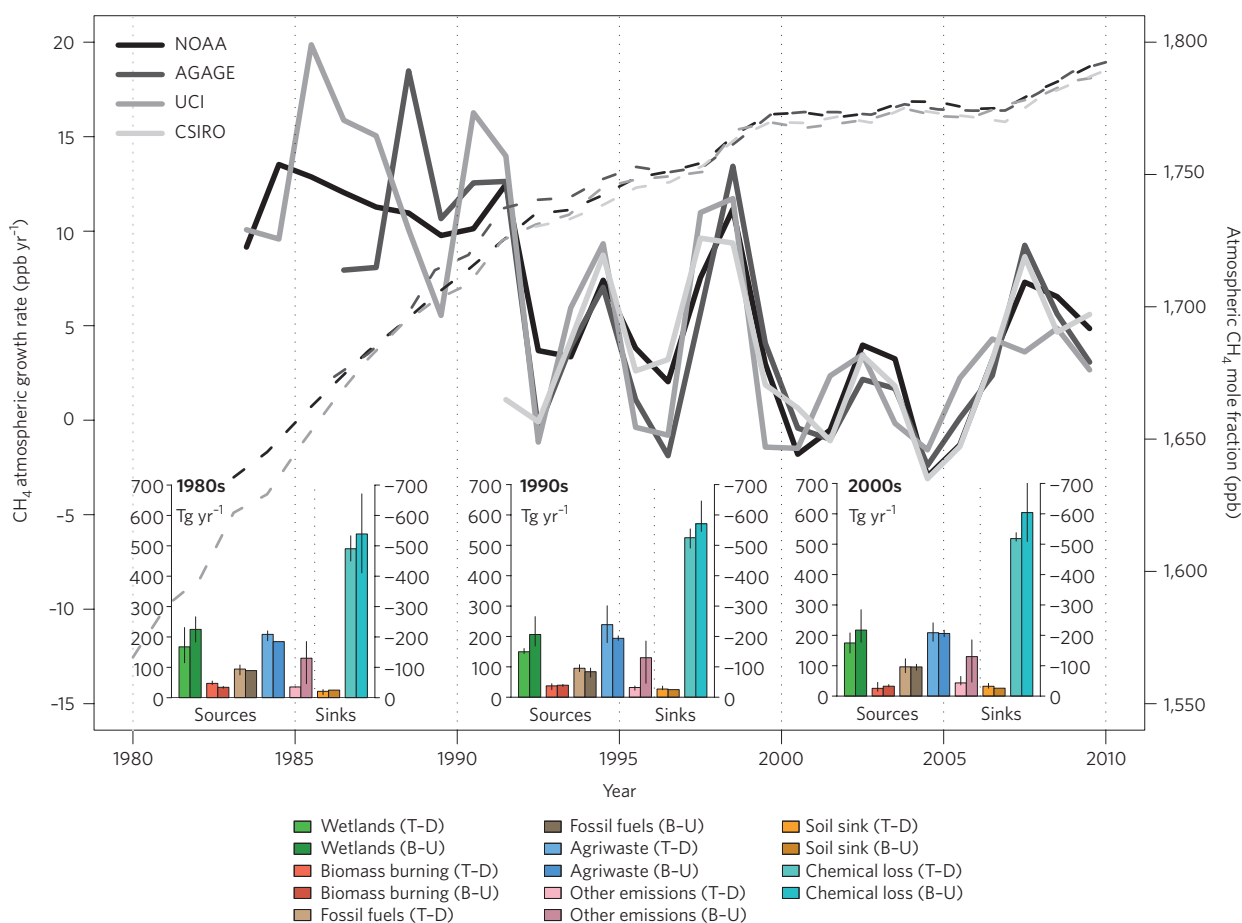


Figure 1 | Evolution of the atmospheric global mole fraction, growth rate and budget of methane for the past three decades. The mole fraction (dashed lines) and growth rate (solid lines) from NOAA, AGAGE, UCI and CSIRO networks are shown in varying shades of black/grey. Bar charts show global decadal surface emissions and sinks calculated from top-down (T-D, light-coloured bars) and bottom-up (B-U, dark-coloured bars) approaches. Categories are split into: natural wetlands, biomass burning, fossil fuels, agriculture and waste, other sources (see Table 1), soil uptake and chemical loss by OH oxidation. Error bars spread between minimum and maximum values.

chlorine radicals and atomic oxygen radicals in the stratosphere¹⁷ (~3%), and reactions with chlorine radicals from sea salt in the marine boundary layer²⁹ (~3%).

Global decadal budget

We combine state-of-the-art top-down and bottom-up approaches (Box 1) using a consistent methodology (see Methods) to assess global CH₄ sources and sinks over the past three decades. At the global scale for the 2000s, top-down inversions yield total global emissions of 548 Tg of CH₄ per year with a minimum–maximum range of 526–569 (six models in Table 1) and a global sink of 540 [514–560] Tg CH₄ yr⁻¹. The source–sink mismatch reflects the observed average imbalance of 6 Tg CH₄ yr⁻¹ of the CH₄ growth rate in the 2000s, which is smaller than that of the 1980s and 1990s (34 Tg CH₄ yr⁻¹ and 17 Tg CH₄ yr⁻¹, respectively; Fig. 1). In fact, stabilization of atmospheric CH₄ prevailed in the early 2000s, and the atmospheric increase resumed after 2006.

Summing up all bottom-up emission estimates, a different picture emerges for the global source for the 2000s. We obtain a value of 678 Tg CH₄ yr⁻¹, which is 20% larger than the inversion-based estimate ($P < 0.01$; Table 1). The higher global source in bottom-up estimates is explained by a larger sum of natural emissions (from wetlands, freshwater, and geological sources) than in the inversions (Table 1). For the 2000s, the bottom-up estimate of the total sink is 632 Tg CH₄ yr⁻¹, with a large range (592–785). Most of this sink — 604 Tg CH₄ yr⁻¹ — is due to the hydroxyl radical CH₄ sink, as estimated

by the nine bottom-up chemistry climate models (CCMs)³⁰. The OH sink simulated by the seven models that run time slices from the 1980s to the 2000s is found to increase with time, which contrasts with the stability of the OH sink inferred from top-down inversions for the 1990s and the 2000s (Table 1). The positive trend in the OH sink in the CCMs can be explained by the fact that the chemical consumption of OH, for instance through reactions with CH₄ and carbon monoxide, is offset by the production of OH through photochemical reactions, involving water vapour, nitrogen oxides and stratospheric ozone. The stable OH sink inferred from top-down inversions relates to the observed atmospheric record of methyl chloroform, which is used to infer OH changes on decadal scales³⁰.

We group decadal estimates of emissions (top-down and bottom-up) into five categories: natural wetlands; other natural emissions (termites, geological, fresh water systems, permafrost and hydrates); agriculture and waste; fossil fuels; and biomass and biofuel burning (Table 1). Freshwater systems include lakes, reservoirs, streams and rivers. In the 2000s, natural wetland emissions (top-down, 142–208 Tg CH₄ yr⁻¹; and bottom-up, 177–284 Tg CH₄ yr⁻¹) and agriculture and waste emissions (top-down, 180–241 Tg CH₄ yr⁻¹; and bottom-up, 187–224 Tg CH₄ yr⁻¹) dominate CH₄ emissions, followed by anthropogenic fossil fuel emissions, other natural emissions and emissions from biomass and biofuel burning (Table 1). Together with natural CH₄ emissions from lake and freshwater sources^{31,32}, we find an imbalance of almost 50 Tg CH₄ yr⁻¹ (in the 2000s) between the mean global emission and the mean global sink in the

bottom-up approach, which is larger than the observed growth rate of around 6 Tg CH₄ yr⁻¹.

This discrepancy, combined with the fact that the global mean emission is 130 Tg CH₄ yr⁻¹ greater in the bottom-up approach than in the top-down approach (Table 1), suggests that CH₄ emissions are overestimated in the bottom-up approach. Indeed, the bottom-up global emission estimate is obtained by adding up independently estimated flux components, and thus lacks a constraint on its global magnitude. In contrast, the global CH₄ emission derived from the top-down approach is constrained at the global scale by the atmospheric CH₄ growth rate, using atmospheric CH₄ measurements, and by the magnitude of the chemical sink, using proxy atmospheric observations, such as the concentration of methyl chloroform, to estimate OH concentrations. Such proxy methods have proven to be reliable indicators of mean OH levels in the troposphere, although their ability to capture OH changes has been widely discussed^{33,34}. These proxy methods suggest that the mean global chemical sink for CH₄ derived from bottom-up estimates may also be overestimated, especially in the 2000s (Table 1).

When summing up anthropogenic fossil emissions, natural fossil CH₄ from onshore and offshore seeps^{35,36} (part of geological emissions in Table 1) and hydrates, bottom-up total fossil emissions account for 28% (~156 Tg CH₄ yr⁻¹) of the global CH₄ source between 1985 and 2000. This is consistent with an analysis of ¹⁴C-CH₄ atmospheric measurements³⁷ in both hemispheres inferring a 30 ± 2% fossil fraction in the global CH₄ source. However, fossil emissions of this magnitude are not confirmed by a recent analysis of the global atmospheric record of ethane¹⁵, which is co-emitted with geological CH₄. Top-down inversions cannot provide useful information to settle this debate, as they generally do not separate this source from other natural emissions (Table 1). Consideration of the natural fossil CH₄ source, neglected in previous Intergovernmental Panel on Climate Change (IPCC) assessments, thus represents a significant update to the global CH₄ budget, although it is still debated.

Global budget uncertainty

Uncertainties associated with decadal CH₄ budgets are expressed by the minimum–maximum range between different decadal estimates, due to the small number of studies available for calculating

reliable standard deviations (Table 1). For the 2000s, the uncertainty range for bottom-up estimates — defined as (max–min)/mean — is 50% for natural wetlands and typically 100% for other natural sources, though the other individual natural sources have smaller fluxes than wetlands. Anthropogenic sources seem to be known more precisely, with an uncertainty range of 30% for agriculture/waste- and fossil-fuel-related emissions, and 20% for biomass burning. The uncertainty range of the global sink is 40%, but drops to 20% when removing one outlier with very high total OH loss in a recent comparison of climate chemistry models^{30,38}. Note that the uncertainties reported in Table 1 are correlated to some extent. Because of more recent and robust estimates for each decade, each term in the budget has a smaller error range than in the IPCC AR4 report: 50% smaller for wetlands, 60% smaller for biomass burning, and 40% smaller for agriculture and waste emissions (Table 1).

Natural wetlands have the largest absolute uncertainty of any of the emission categories, with a min–max range of 107 Tg CH₄ yr⁻¹ in the bottom-up approach (177–284 Tg CH₄ yr⁻¹). This large range is confirmed by a recent multi-model analysis³⁹ showing a ±40% range of wetland emissions around an average of 190 Tg CH₄ yr⁻¹. In the three wetland emission models used here^{40–42}, emissions were calculated for each grid point as the product of a flux rate and a wetland area, both having uncertainties. Uncertainties in wetland extent seem to be the dominant source of discrepancy in modelled CH₄ emissions^{39,43}.

The OH sink seems to have a smaller error range using proxy methods in the top-down approach (max–min range of 30 Tg CH₄) than in bottom-up CCMs (max–min range of 250 Tg CH₄, dropping to 110 Tg CH₄ when removing one outlier model from the Atmospheric Chemistry and Climate Model Intercomparison Project (ACCMIP)^{30,38}), in which different humidity and temperature fields cause a large spread of the OH sink³⁸.

Following IPCC AR5 guidelines for the treatment of uncertainties⁴⁴, we defined a level of confidence for both top-down estimates and bottom-up estimates, based on robustness (number of published studies) and agreement (difference between maximum and minimum estimates, relative to the mean). Many studies have focused on constraining the CH₄ budget during the 1990s and 2000s, but fewer estimates are available for the 1980s. As a result,

Box 1 | New data to assess the CH₄ budget

The top-down approach is based on atmospheric inversion models, which determine ‘optimal’ surface fluxes^{92,93} that best fit atmospheric CH₄ observations given an atmospheric transport model including chemistry, prior estimates of fluxes, and their uncertainties. Global atmospheric inversions provide a time-varying distribution of CH₄ fluxes, albeit with limited insight into the underlying processes when different sources overlap in the same region. This is, for example, often the case for agricultural, waste and fossil emissions in densely populated areas of east Asia, Europe and North America. We collected results from nine inversion systems (Supplementary Table S1).

The bottom-up approach includes process-based models estimating CH₄ emissions, and CCMs estimating the OH sink. Eight bottom-up models for wetland and fire CH₄ emissions are parameterized with empirical knowledge of local processes and driven by global data sets of climate, or satellite-observed burned area, to simulate CH₄ fluxes on spatial and temporal scales relevant for regional and global budgets (Supplementary Section II). Bottom-up emission inventories^{56,81,82} based on energy use, agricultural activity, and emission factors from different sectors provide yearly or decadal mean estimates of anthropogenic waste-related, rice, livestock, biofuel, and fossil fuel emissions, usually at

national scales. Three inventories for anthropogenic emissions are used, updated to 2008 (Supplementary Information).

The photochemical sink of CH₄ is large and difficult to quantify, given the very short lifetime of OH (~1 sec) and its control by a myriad of precursor species. Direct measurements of atmospheric OH radicals do not have the required accuracy and coverage to derive global OH concentrations and consequently the magnitude of the CH₄ sink. We estimated CH₄ loss due to OH from the output of nine numerical CCMs⁶⁵, which are categorized here as an atmospheric bottom-up approach. The OH concentration as calculated by CCMs can be further adjusted, at a large scale, by inversions based on measurements of tracers with known emissions and whose dominant sink is oxidation by OH, such as methyl chloroform^{34,49,85,94} or chloromethanes^{33,34}.

Combining top-down and bottom-up approaches allows us to investigate the consistency of each term of the CH₄ budget²¹. In this comparison, it should be noted that bottom-up models and inventories are not independent from inversions, because they are usually used in inversions to prescribe a prior spatial, and sometimes temporal, distribution of the emissions and sinks. However, inversions use independent atmospheric observations to partially correct the prior values.

Table 1 | CH₄ budget for the past three decades.

	Tg CH ₄ yr ⁻¹					
	1980-1989		1990-1999		2000-2009	
	Top-down	Bottom-up	Top-down	Bottom-up	Top-down	Bottom-up
Natural sources	203 [150-267]	355 [244-466]	182 [167-197]	336 [230-465]	218 [179-273]	347 [238-484]
Natural wetlands	167 [115-231] ^{19,21,76}	225 [183-266] ^{40,41}	150 [144-160] ^{21,74,77}	206 [169-265] ⁴⁰⁻⁴²	175 [142-208] ^{46,53,73,75,77,86}	217 [177-284] ⁴⁰⁻⁴²
Other sources	36 [35-36] ^{19,21,76}	130 [61-200]	32 [23-37] ^{21,74,77}	130 [61-200]	43 [37-65] ^{46,53,73,75,77}	130 [61-200]
Fresh water (lakes and rivers)		40 [8-73] ^{31,32}		40 [8-73] ^{31,32}		40 [8-73] ^{31,32}
Wild animals		15 [15-15] ¹⁶		15 [15-15] ¹⁶		15 [15-15] ¹⁶
Wildfires		3 [1-3] ^{16,47,55,88,89}		3 [1-5] ^{16,47,55,88,89}		3 [1-5] ^{16,47,55,88,89}
Termites		11 [2-11] ^{16,48,55,91}		11 [2-22] ^{16,37,87,91}		11 [2-22] ^{16,37,87,91}
Geological (incl. oceans)		54 [33-75] ^{35,55,90}		54 [33-75] ^{35,55,90}		54 [33-75] ^{35,55,90}
Hydrates		6 [2-9] ^{16,36,87}		6 [2-9] ^{16,36,87}		6 [2-9] ^{16,36,87}
Permafrost (excl. lakes and wetland)		1 [0-1] ⁵⁵		1 [0-1] ⁵⁵		1 [0-1] ⁵⁵
Anthropogenic sources	348 [305-383]	308 [292-323]	372 [290-453]	313 [281-347]	335 [273-409]	331 [304-368]
Agriculture and waste	208 [187-220] ^{19,21,76}	185 [172-197] ⁵⁶	239 [180-301] ^{21,74,77}	188 [177-196] ^{55,56,81}	209 [180-241] ^{46,53,73,75,77}	200 [187-224] ^{55,56,81}
Biomass burning (incl. biofuels)	46 [43-55] ^{19,21,76}	34 [31-37] ^{78,80}	38 [26-45] ^{21,74,77}	42 [38-45] ^{78,80}	30 [24-45] ^{47,53,72,73,75,77}	35 [32-39] ^{47,78,80,89}
Fossil fuels	94 [75-108] ^{19,21,76}	89 [89-89] ⁵⁶	95 [84-107] ^{21,74,77}	84 [66-96] ^{55,56,81}	96 [77-123] ^{46,53,73,75,77}	96 [85-105] ^{55,56,81}
Sinks						
Soils	21 [10-27] ^{19,21,76}	28 [9-47] ^{27,42}	27 [27-27] ²¹	28 [9-47] ^{27,42,89}	32 [26-42] ^{46,53,73,75,86}	28 [9-47] ^{27,42,89}
Total chemical loss	490 [450-533] ^{19,21,76}	539 [411-671] ^{21,29,38,83}	525 [491-554] ^{21,83}	571 [521-621] ^{21,29,38,83}	518 [510-538] ^{46,53,73,75,77}	604 [483-738] ^{21,29,38,83}
Tropospheric OH		468 [382-567] ^{30,38}		479 [457-501] ^{30,38}		528 [454-617] ^{30,38}
Stratospheric loss		46 [16-67] ^{22,38,83}		67 [51-83] ^{21,38,83}		51 [16-84] ^{21,38,83}
Tropospheric Cl		25 [13-37] ²⁹		25 [13-37] ²⁹		25 [13-37] ²⁹
TOTALS						
Sum of sources	551 [500-592]	663 [536-789]	554 [529-596]	649 [511-812]	548 [526-569]	678 [542-852]
Sum of sinks	511 [460-559]	539 [420-718]	542 [518-579]	596 [530-668]	540 [514-560]	632 [592-785]
Imbalance (sources-sinks)	30 [16-40]		12 [7-17]		8 [-4-19]	
Atmospheric growth rate	34		17		6	

Top-down and bottom-up estimates are listed separately for the different categories in Fig. 1. For top-down inversions, the 1980s decade starts in 1984. Numbers in square brackets represent minimum and maximum values. A balance with the atmospheric annual increase and the sum of the sources has been assumed for inversions not reporting their global sink. Stratospheric loss for bottom-up is the sum of the loss by radicals, a 10 Tg yr⁻¹ loss due to O(1D) radicals²² and a 20-35% contribution due to Cl radicals²⁹. Ranges of total chemical loss are about half the reported ranges (for example, [509-619] for the 2000s) when removing one outlier.

estimates for all source categories during the 2000s are more robust, especially for inversions (Fig. 2). Agreement among studies is high (difference is less than 33%) for agriculture and waste (top-down and bottom-up), biomass burning and fossil fuels (bottom-up) and OH loss (top-down), whereas agreement is only medium (33-66% difference) for natural wetlands (top-down and bottom-up), fossil fuel emissions (top-down) and OH sink (bottom-up) estimates. Low agreement (> 66% difference) is found for biomass burning (top-down) and other natural sources (bottom-up). Increasing the number of studies does not necessarily lead to enhanced agreement.

This can be seen for the fossil fuel and other sources categories, partly because of poorly constrained models, and partly because the results from a single new study can produce a large increase in the spread of emission estimates when very few studies are available.

No source or sink category reaches the highest level of confidence (highest agreement and highest robustness), emphasizing the large uncertainties that remain in our understanding of CH₄ emissions. Overall, higher confidence in global emissions is found for agriculture and waste (top-down) than for fossil fuels, the OH sink, natural wetlands and other natural sources.

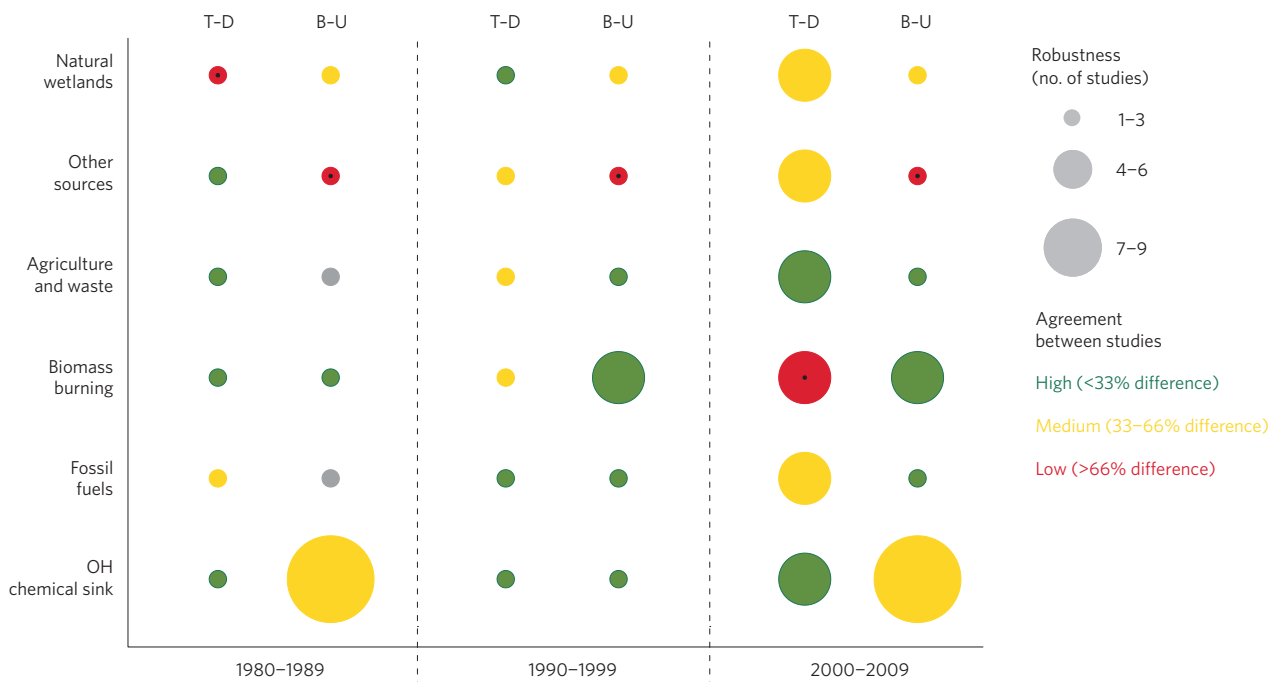


Figure 2 | Evolution of uncertainty on estimates of methane emissions and sinks presented in Table 1. Circle size depicts the robustness of the estimate (number of studies). Circle colour illustrates the level of agreement among studies (min-max ranges): green, high confidence; yellow, medium confidence; red (with black dot), low confidence. Circles are grey when only one study has been used. A large green circle, for example, indicates a very good level of confidence⁴⁴.

Regional decadal budget

The geographical breakdown of emissions per category and per region reveals major CH₄ emission zones worldwide and the level of consistency between top-down and bottom-up approaches (Fig. 3 and Supplementary Section ST2 and Tables S2 and S3). Anthropogenic emissions dominate in Europe, North America, China, and the fossil-fuel-producing countries of eastern Europe and central Asia, with good agreement between top-down and bottom-up approaches (Fig. 3). Emission ranges are given in Table S2. Densely populated regions usually emit fossil, agricultural and waste CH₄, making these sources difficult to separate in top-down inversions. Noteworthy is the large range of estimates for anthropogenic fossil CH₄ emissions from China in the top-down approach, possibly due to the low density of atmospheric CH₄ measurements in this region, and to biases in inventories⁴⁵. The large range of anthropogenic CH₄ emission estimates in Europe and North America possibly reflects uncertainties in emission factors, and in the partition between waste and fossil CH₄ sources. In emerging economies, agriculture and waste emissions are highest in China (top-down, 29 Tg CH₄ yr⁻¹; bottom-up, 28 Tg CH₄ yr⁻¹) and India (top-down, 27 Tg CH₄ yr⁻¹; bottom-up, 22 Tg CH₄ yr⁻¹), but are also important in southeast Asia and temperate South America due to extensive rice agriculture and livestock industries (Supplementary Table S2). In India and China, agriculture and waste constitutes the single largest regional source of CH₄. However, per capita CH₄ emissions in India and China are still 35% and 85%, respectively, of the mean for OECD countries.

When aggregated over large regions, wetlands dominate emissions in tropical South America (top-down, 28 Tg CH₄ yr⁻¹; bottom-up, 58 Tg CH₄ yr⁻¹) and Africa (top-down, 36 Tg CH₄ yr⁻¹; bottom-up, 24 Tg CH₄ yr⁻¹), with significant emissions in southeast Asia, temperate South America, boreal North America and boreal Eurasia (Supplementary Table S2). Tropical South America shows the largest regional discrepancy between top-down (17–48 Tg CH₄ yr⁻¹) and bottom-up (39–92 Tg CH₄ yr⁻¹) wetland emissions (Supplementary Tables S2 and S3). The seven inversions

using only surface measurements give the lowest estimates for the 2000s decadal mean wetland emission (17–30 Tg CH₄ yr⁻¹), and the two inversions using SCIAMACHY column satellite data combined with surface measurements⁴⁶ (27 and 48 Tg CH₄ yr⁻¹) agree better with bottom-up estimates (39–92 Tg CH₄ yr⁻¹). Only short time series of CH₄ *in situ* measurements are available for inland South America, which makes it one of the least constrained regions for inversions using surface measurements. The wetland models used in this study simulate large emissions in the Amazon region, equatorial tropical Africa, tropical Asia (for example, Bangladesh, India, China and Indonesia), Canada and boreal Eurasia. Simulated emission areas are consistent between models for 66 ± 9% of global wetland emissions over the period 1990–2006 (Supplementary Fig. S0).

When aggregated over large regions, emissions from biomass burning are the largest in Africa (top-down, 9 Tg CH₄ yr⁻¹; bottom-up, 8 Tg CH₄ yr⁻¹) and in tropical South America (top-down, 5 Tg CH₄ yr⁻¹; bottom-up, 4 Tg CH₄ yr⁻¹), but play only a minor role in temperate and boreal regional budgets. The bottom-up estimates are likely to be conservative compared to top-down estimates, as small fires are often undetected by satellite retrieval algorithms⁴⁷. For biomass burning, simulated emission areas are consistent between models for 38 ± 9% of global emissions over the period 1997–2000, revealing robust large emission zones around the thermal equator in Africa (for example, Central African Republic, Democratic Republic of the Congo, Republic of the Congo, Angola, Zambia and Cameroon), central South America (Brazil and Bolivia), Indonesia, and to a lesser extent in eastern Russia, Laos, and Mexico (Supplementary Fig. S0). Emission zones in northern Australia and in boreal regions (Canada and Siberia) can also be clearly identified.

Other natural sources, including termites, lakes and other fresh waters, and onshore geological emissions show maximum values in Africa and tropical South America, due to the relatively strong contribution of emissions by termites⁴⁸. A new empirical model of termite CH₄ emissions developed in this study indicates that Africa and tropical South America are major contributors to the global

termite source, contributing 30% and 36%, respectively, of the total (Supplementary Section ST7). Finally, CH₄ loss due to OH radicals is largest in the tropical atmosphere, both over land and oceans, as the tropics are the major region of OH production⁴⁹.

Attribution of temporal changes

Year-to-year variations of CH₄ fluxes have been intensively studied^{4,14,21,47,50}. The present study confirms the findings from previous ones showing that, over the last three decades, variations in wetland emissions have dominated the year-to-year variability in surface emissions (Supplementary Fig. S5). Interannual variability in wetland emissions surpasses that of biomass burning emissions, except during intensive fire periods^{21,50}. Analyses of anomalies in CH₄ fluxes following the Mount Pinatubo^{21,51} eruption in 1991 and the record-high El Niño^{47,52} in 1997–1998 are summarized in Supplementary Sections ST4 and ST5. Both models and observations compiled in the present study consistently describe small interannual variability in the OH sink in the 2000s compared with the previous two decades (<3%, 1σ of annual means; Supplementary Section ST6), in line with previously reported estimates (<5%)^{34,53}.

The observed decadal changes remain much more enigmatic than yearly anomalies (Supplementary Fig. S5). We use a scenario approach, built from our synthesis and from recent publications, to investigate these changes, and the contribution of the different CH₄

sources to them (see Methods). We assume that decadal changes in global mean CH₄ emissions since 1985 are well represented by the mean of those five atmospheric inversions covering the past three decades⁵³, averaged on a five-year basis (Fig. 4 and Methods). A global mass balance model⁵⁴ based on the atmospheric observations of the four surface networks and on possible changes in CH₄ lifetime is used to provide uncertainties on the mean inversion (blue shaded area at the top of Fig. 4). These observation-driven global CH₄ emissions show three distinct regimes: an increase before 1990, an oscillation around a constant mean value during 1990–2005, and an increase after 2006^{4,14,53}. A storyline (S₀) is constructed by adding wetland emissions from top-down inversions (average of five inversions) to other estimates (EPA (ref. 55) and EDGARv4.2 (ref. 56) inventories).

1985–2005. The S₀ storyline clearly overestimates global emissions after 1990, which calls for corrections to the magnitude of one or several sources in the S₀ scenario (Fig. 4). Using ethane firn air and atmospheric measurements, two recent studies indicated that CH₄ emissions from the fossil fuel sector decreased between 1985 and 2000 at a rate of –0.4 to –0.8 Tg CH₄ yr^{–1}, and attributed such a decline to decreasing fugitive emissions (leaks during extraction, treatment and use of fossil fuels) from oil and gas industries^{15,57}. One of these studies further extended the ethane record up to 2010¹⁵, with either a slower decline or a stabilization of fossil fuel emissions

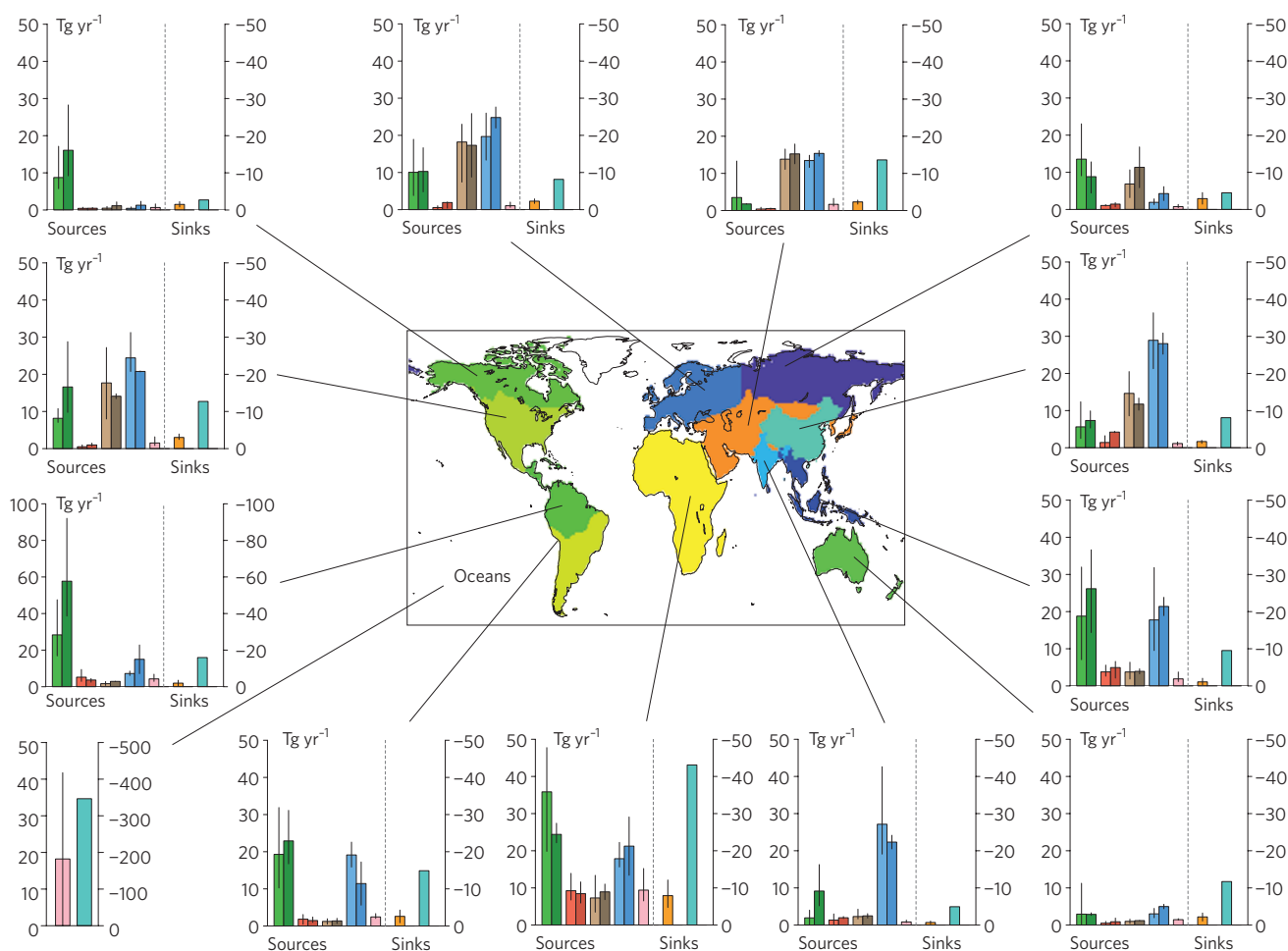


Figure 3 | Regional budgets for 2000–2009 over 13 regions. The considered regions are nine TransCom regions⁸⁴, plus separate regions for India, China and southeast Asia, and one region for oceans. Source and sink categories are the same as in Fig. 1. Both top-down (T-D, light-coloured bars) and bottom-up (B-U, dark-coloured bars) approaches are shown. Oceans are considered as one large region (bar chart at the bottom left), with ocean emissions (pink) and chemical loss over the ocean (turquoise). Error bars indicate the spread between the minimum and the maximum values.

after 2000 (see Fig. 4 of ref. 15 and Methods). Indeed, an intensified coal exploitation^{45,56} after 2000 may have offset a decline in fugitive emissions. In parallel, rice paddy emissions have decreased (~ -0.4 to -0.8 Tg CH₄ yr⁻¹) according to the EDGAR4.2 inventory⁵⁶ during the 1980–2000 period, and remained stable between 2000 and 2005. Assuming that CH₄ fossil fuel fugitive emissions decreased between 1985 and 2000⁵³ and were stable from 2000 to 2005¹⁵, and keeping the other sources as in S₀, leads to a first plausible scenario that is consistent with the observation-driven global emissions (S₁ in Fig. 4). An alternative scenario (S'₁), using bottom-up ecosystem model results for wetland emissions as a storyline instead of top-down inversions, is also consistent with the observation-driven global emissions.

Two different analyses of $\delta^{13}\text{C}$ -CH₄ isotopic composition trends^{58,59} for 1990–2005 reached contradictory conclusions. In one, constant fossil fuel emissions but decreasing microbial emissions in the Northern Hemisphere were inferred⁵⁸, the latter mainly attributed to decreasing rice emissions. In the other⁵⁹, fossil fuel and microbial emissions remained constant. Assuming constant fossil fuel emissions during 1985–2005 and decreasing microbial emissions⁵⁸ produces a second scenario that is mostly consistent with observation-driven global emissions when using wetland fluxes from top-down inversions (S₂ in Fig. 4), but not when using wetland fluxes from bottom-up ecosystem models (S'₂). Assuming decreasing fossil fuel emissions before 1990 (as in S₁), but constant fossil fuel and microbial emissions between 1990 and 2005⁵⁹, produces a third scenario that is consistent with observation-driven global emissions, with either top-down or bottom-up wetland emission estimates (S₃ and S'₃ in Fig. 4).

Overall, the three plausible scenarios, among many other possible source compositions matching global decadal changes, suggest that a decrease in fossil fuel CH₄ emissions is a more likely explanation for the stability of global CH₄ emissions between 1990 and 2005 than a reduction in microbial CH₄ emissions. An actual decrease in rice paddy emissions may have been surpassed by an increase in other microbial emissions (natural wetlands, animals, landfills and waste) as found by ecosystem models combined with the EDGAR4.2 inventory. Considering the significant uncertainties reported in a recent isotope study⁵⁹ for the 1990–2005 period, decreasing-to-stable fossil fuel emissions, combined with stable-to-increasing total microbial emissions, would reconcile the atmospheric ethane trends with the ¹³C-CH₄ trends, at least for one ¹³C-CH₄ data set⁵⁹. Finally, trends in the magnitude of the OH CH₄ sink, which remain uncertain over decadal timescales, can still modulate these incomplete conclusions³⁴.

The increase resumes from 2006 onwards. Atmospheric CH₄ levels resumed growth after 2006¹⁴, with inferred global emissions being 17–22 Tg CH₄ yr⁻¹ greater around 2010 than around 2005 (five-year basis averages; top of Fig. 4). Several studies concluded that a recent surge in natural wetland emissions is one main cause of increasing CH₄ levels, in response to abnormally high temperatures in northern high latitudes in 2007, and increased rainfall over tropical wetlands during 2008–2009 and 2010–2011^{13,53,60}, two La Niña periods⁴. Furthermore, fossil fuel CH₄ emissions probably increased again after 2005, mostly due to the intensification of shale gas and oil extraction in the United States and coal exploitation by the Chinese and Indian economies⁴⁵.

After 2005, the three scenarios use fossil fuel emission changes from the EPA inventory, and the average of EPA and EDGAR4.2 inventories for all other sources barring natural wetlands. Microbial and fossil fuel sources for all scenarios show positive trends after 2005, resulting in an increase of global emissions of 23–33 Tg CH₄ yr⁻¹ around 2010 as compared to around 2005 (five-year basis averages). This is a 30% overestimation compared with the mean increase derived from the observations

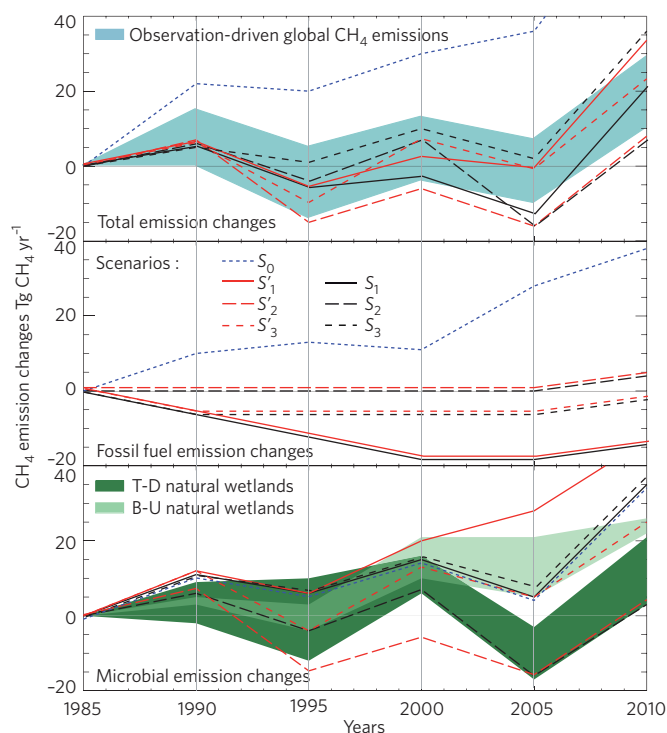


Figure 4 | Plausible scenarios explaining changes in methane emissions over the past three decades.

Different lines depict different scenarios of five-year-averaged emission changes since 1985 (see Methods): S₀ (dotted blue lines), S₁ and S'₁ (solid black and red lines), S₂ and S'₂ (long-dashed black and red lines), S₃ and S'₃ (short-dashed black and red lines) around a mean inversion (Methods and Supplementary Section ST5). Top: range of global CH₄ emission changes (blue shaded area) and total emission changes (Methods and Supplementary Section ST5). Middle: emission changes from fossil fuels (coal, gas and oil industries). Bottom: emission changes from microbial sources (natural wetlands, rice, animals and waste). The dark and light green shaded areas represent the range of top-down (T-D) and bottom-up (B-U) model results, respectively, for natural wetland emissions.

(17–22 Tg CH₄ yr⁻¹, see above). Thus, either the increase in fossil fuel emissions is overestimated by inventories, or the sensitivity of wetland emissions to precipitation and temperature is too large in some wetland emission models³⁹. The contribution of microbial versus fossil emissions to this increase remains largely uncertain; respective contributions vary from 20 to 80%, if accounting for all additional top-down inversions available for the 2000s (Supplementary Fig. S5 and Table 1).

Shortcomings and uncertainty reductions

Our analyses suggest four main shortcomings in the assessment of regional to global CH₄ budgets. First, decadal means and interannual changes in CH₄ emissions from natural wetlands and freshwater systems are too uncertain. It is critically important to improve wetland mapping, both by refining land surface models (for example, through improving estimates of tropical flood plains in hydrological models, specific model developments for peatlands, and the integration of freshwater systems) and by further developing remotely sensed inundation data sets⁶¹ (for instance for dense tropical forests). The scarcity of wetland CH₄ flux measurements and data sets limits the ability to validate large-scale modelled CH₄ emissions for natural wetlands and fresh waters⁴³. The extension of the CO₂ FLUXNET measurements and database⁶² to CH₄ fluxes is probably achievable at a reasonable cost, and would provide useful constraints for land surface models. For interannual variations in wetland emissions, the sensitivity of emission rates to warming at

high northern latitudes and to rainfall changes in the tropics needs to be more consistently quantified in wetland models. The Amazon drought in 2010⁶³ should have resulted in a drop in wetland CH₄ emissions, and ongoing analyses may allow researchers to test the hypothesis that tropical wetland CH₄ emissions respond strongly to rainfall anomalies and trends.

Second, the partitioning of CH₄ emissions by region and process is not sufficiently constrained by atmospheric observations in top-down models. Regional partitioning of total emissions would benefit from denser and more evenly distributed CH₄ concentration data. This can be achieved by further developing synergies between high precision monitoring of the surface and the lower atmosphere, including poorly sampled key areas such as the Amazon Basin, Siberia and tropical Africa on one hand, and retrievals of global-scale CH₄ columns by satellites and by high precision remote sensing from the ground on the other. Including continuous measurements of the δ¹³C stable isotope (¹³CH₄) at surface stations would help separate biogenic emissions from other sources. Measurements of the δD stable isotope (CH₃D) would provide constraints on the uncertain OH CH₄ sink, which can also be constrained by new proxy tracers^{33,34}. Radiocarbon CH₄ data (¹⁴CH₄) would help constrain the uncertain fossil part of CH₄ emissions, if ¹⁴CH₄ emissions from nuclear installations can be accurately estimated³⁷. Estimating long-term trends of fluxes and concentrations requires equally long-term observations, which in turn require stable and coordinated networks⁶⁴.

Third, decadal trends in natural and anthropogenic emissions are still very uncertain and limit our ability to definitively attribute changes in emissions from specific sources to observed atmospheric changes since the 1990s. In addition to the (already noted) improvements in land surface models required, inventories for anthropogenic emissions should systematically include an uncertainty assessment, and should improve their representation of emission trends (for instance by more frequently updating the time-dependent factors used in their calculations).

Fourth, uncertainties in the modelling of atmospheric transport and chemistry limit the optimal assimilation of atmospheric observations by increasing uncertainties in top-down inversions. Such uncertainties are also only partly estimated in current inversions. We therefore recommend the continuation of ongoing international model inter-comparisons, which can provide a quantification of transport and chemistry errors to be included in top-down inversions^{65,66}.

From challenge to opportunity

Our decadal CH₄ budgets reveal that bottom-up models may overestimate total natural CH₄ emissions. The various emission scenarios tested — designed to explain the temporal changes in atmospheric CH₄ levels observed in this and previous studies — suggest that the stabilization of atmospheric CH₄ in the early 2000s is likely to be due to a reduction in or stabilization of fossil fuel emissions, combined with a stabilization of or increase in microbial emissions. After 2006, the renewed global increase in atmospheric CH₄ is consistent with higher emissions from wetlands and fossil fuel burning, but the relative contributions remain uncertain.

In the context of climate change mitigation, atmospheric CH₄ poses both an opportunity and a challenge. The challenge lies in more accurately quantifying the CH₄ budget and its variations. Our synthesis suggests that improvements in models of natural wetland and freshwater emissions, the integration of surface networks monitoring CH₄ concentrations and fluxes (including isotopic composition) and new satellite missions (including active space-borne observations⁶⁷), improvements in anthropogenic emission trends in inventories, and uncertainty reductions in models of atmospheric transport and chemistry, could all help. The opportunity lies in the possibility of developing short-term

climate change mitigation policies that take advantage of the relatively short atmospheric lifetime of CH₄ of about 10 years, and the known technological and agronomical options available for reducing emissions⁶⁸.

The potential intensive exploitation of natural gas from shale formations around the world may lead to significant additional CH₄ release into the atmosphere⁶⁹, although the potential magnitude of these emissions is still debated⁷⁰. Such additional emissions, and combustion of this 'new' fossil fuel source, may offset mitigation efforts and accelerate climate change. In the longer term, the thawing of permafrost or hydrates could increase CH₄ emissions significantly, and introduce large positive feedbacks to long-term climate change⁷¹. A better quantification of the global CH₄ budget, with regular updates as done for carbon dioxide⁷², will be key to both embracing the opportunities and meeting the challenge.

Methods

Data analysis. Top-down and bottom-up studies addressing the evolution of the CH₄ cycle after 1980 and covering at least five years of a decade were gathered. Therefore, the number and the nature of studies used in this work vary from one decade to another. Top-down inversions include atmospheric chemistry transport models and assimilation systems^{19,46,53,73–77}. Bottom-up approaches comprise modelling studies for wetland^{40–42} and biomass-burning emissions^{47,78–80}, emission inventories for anthropogenic^{55,56,81} and natural sources⁸², and a suite of atmospheric chemistry models within the ACCMIP intercomparison project providing CH₄ chemical loss^{30,39,83}.

The monthly fluxes (emissions and sinks) provided by the different groups were post-processed similarly. They were re-gridded on a common grid (1° × 1) and converted into the same units (Tg CH₄ per grid cell); then monthly, annual and decadal means were computed for 12 regions based on the TransCom⁸⁴ intercomparison map, with subdivisions in high-emission regions. Regional and global means were used to construct Figs 1, 2 and 3, Supplementary Figs S2 and S3, Table 1 and Supplementary Tables S2 and S3.

The reported ranges and error bars represent the minimum and maximum values obtained among the different studies (Figs 1, 3 and 4 and Table 1). The small number of studies for some categories makes it difficult to properly apply a standard deviation.

Interannual variability (IAV) was computed as the difference between the 12-month running mean and the long-term mean. However a consistent period for estimating the long-term mean was not compatible with all data sources (Supplementary Fig. S5).

Observation-driven global CH₄ emissions. For 'attribution of temporal changes', we used the only top-down study that estimates CH₄ emissions over the past 30 years⁵³ with five different set-ups. The mean of these five inversions was assumed to represent average global emissions. However these five inversions only partially represented the full range of global CH₄ emissions, due to differences in prior emission scenarios and errors, observations and their errors, OH fields and atmospheric transport representation. To estimate the full range of global CH₄ emissions we complemented the mean inversion with a sensitivity analysis based on a one-box model for the whole atmosphere³⁴. The change in the global burden of CH₄ is given by:

$$\frac{d[\text{CH}_4]}{dt} = E - \frac{[\text{CH}_4]}{\tau} \quad (1)$$

where [CH₄] is the global CH₄ burden, E is the sum of all emissions, and τ is the total atmospheric CH₄ lifetime. Equation (1) can be rearranged to calculate the annual CH₄ source strength E as follows:

$$E = \frac{d[\text{CH}_4]}{dt} + \frac{[\text{CH}_4]}{\tau} \quad (2)$$

In this equation, the annual increase d[CH₄]/dt and the burden [CH₄] were given by the yearly-averaged growth rates and mole fractions of Fig. 1. Global CH₄ emissions were generated by computing emissions with equation (2) for each of the four networks and for a lifetime τ varying from 8 to 10 years to include uncertainties in OH changes^{34,85}. Minimum and maximum values of E were extracted for five-year periods to produce the range of emissions plotted around the mean of atmospheric inversions (blue shaded area in Fig. 4, top panel).

Emission scenarios. The emission scenarios are based on five-year average CH₄ fluxes around the years 1985, 1990, 1995, 2000, and 2005. For 2010 we used available years between 2008 and 2012, mainly before 2010. Flux changes from 2005 to 2010 might be slightly biased by missing years after 2010. For example, fossil and microbial emissions both increase between 2005 and 2009; if after 2010 these emissions were further increasing (or decreasing), then the 2005–2010 changes will be underestimated (or overestimated). We assume that such a potential bias does not modify the (mostly) qualitative message of our scenario analysis. The five-year

changes from biomass burning remain small (<2 Tg CH₄ per five-year period) and were not considered here.

The scenarios presented in Fig. 4 use either natural wetland emissions from top-down inversions (S₀) or bottom-up models (S'₀). Other data are taken from recent publications and EDGAR4.2 and EPA inventories.

S₀ and S'₀ are built by summing the mean wetland emissions from inversions and the mean of EPA (ref. 55) and EDGAR4.2 (ref. 56) 'other' emissions. Scenarios S₁ and S'₁ sum the mean wetland emissions with decreasing (1985–2000), constant (2000–2005), and increasing (2005–2010) fossil fuel emissions to be compatible with a recent analysis¹⁵. Scenarios S₂ and S'₂ sum the mean wetland emissions with constant (1985–2005) and increasing (2005–2010) fossil fuel emissions. Other microbial emissions (mean of EPA and EDGAR) are scaled to a recent study⁵⁸. Scenarios S₃ and S'₃ sum the mean wetland emissions with decreasing (1985–1990), constant (1990–2005), and increasing (2005–2010) fossil fuel emissions. Other microbial emissions (mean of EPA and EDGAR) are scaled to remain constant during 1990–2005 according to another recent study⁵⁹. After 2005, all scenarios include fossil fuel emission changes from the EPA inventory, wetland emission changes from inversions or bottom-up studies and other emission changes from the mean of EPA and EDGAR4.2.

Received 17 April 2012; accepted 22 August 2013;
published online 22 September 2013

References

- Etheridge, D. M., Pearman, G. I. & Fraser, P. J. Changes in tropospheric methane between 1841 and 1978 from a high accumulation-rate Antarctic ice core. *Tellus* **44B**, 282–294 (1992).
- Blake, D. R. *et al.* Global increase in atmospheric methane concentrations between 1978 and 1980. *Geophys. Res. Lett.* **9**, 477–480 (1982).
- Cunnold, D. M. *et al.* *In situ* measurements of atmospheric methane at GAGE/AGAGE sites during 1985–2000 and resulting source inferences. *J. Geophys. Res.: Atmos.* <http://dx.doi.org/10.1029/2001jd001226> (2002).
- Dlugokencky, E. J. *et al.* Observational constraints on recent increases in the atmospheric CH₄ burden. *Geophys. Res. Lett.* **36**, L18803 (2009).
- Francey, R. J., Steele, L. P., Langenfelds, R. L. & Pak, B. C. High precision long-term monitoring of radiatively active and related trace gases at surface sites and from aircraft in the southern hemisphere atmosphere. *J. Atmos. Sci.* **56**, 279–285 (1999).
- World Data Centre for Greenhouse Gases (WMO/WDCGG) (2012); <http://ds.data.jma.go.jp/gmd/wdcgg/introduction.html>.
- Brenninkmeijer, C. A. M. *et al.* Civil Aircraft for the regular investigation of the atmosphere based on an instrumented container: The new CARIBIC system. *Atmos. Chem. Phys.* **7**, 4953–4976 (2007).
- Wecht, K. J. *et al.* Validation of TES methane with HIPPO aircraft observations: implications for inverse modeling of methane sources. *Atmos. Chem. Phys.* **12**, 1823–1832 (2012).
- Schuck, T. J. *et al.* Distribution of methane in the tropical upper troposphere measured by CARIBIC and CONTRAIL aircraft. *J. Geophys. Res.: Atmos.* **117**, D19304 (2012).
- Crevoisier, C. *et al.* Tropospheric methane in the tropics — first year from IASI hyperspectral infrared observations. *Atmos. Chem. Phys.* **9**, 6337–6350 (2009).
- Frankenberg, C. *et al.* Global column-averaged methane mixing ratios from 2003 to 2009 as derived from SCIAMACHY: Trends and variability. *J. Geophys. Res.: Atmos.* **116**, D04302 (2011).
- Morino, I. *et al.* Preliminary validation of column-averaged volume mixing ratios of carbon dioxide and methane retrieved from GOSAT short-wavelength infrared spectra. *Atmos. Meas. Tech.* **4**, 1061–1076 (2011).
- Dlugokencky, E. J., Nisbet, E. G., Fisher, R. & Lowry, D. Global atmospheric methane: budget, changes and dangers. *Phil. Trans. R. Soc. A* **369**, 2058–2072 (2011).
- Rigby, M. *et al.* Renewed growth of atmospheric methane. *Geophys. Res. Lett.* **35**, L22805 (2008).
- Simpson, I. J. *et al.* Long-term decline of global atmospheric ethane concentrations and implications for methane. *Nature* **488**, 490–494 (2012).
- Denman, K. L. *et al.* in *IPCC Climate Change 2007: Couplings Between Changes in the Climate System and Biogeochemistry* (eds Solomon, S. *et al.*) (Cambridge Univ. Press; 2007).
- Cicerone, R. J. & Oremland, R. S. Biogeochemical aspects of atmospheric methane. *Glob. Biogeochem. Cycles* **2**, 299–327 (1988).
- Ehhalt, D. H. The atmospheric cycle of methane. *Tellus* **26**, 58–70 (1974).
- Fung, I. *et al.* Three-dimensional model synthesis of global methane cycle. *J. Geophys. Res.* **96**, 13033–13065 (1991).
- Monteil, G. *et al.* Interpreting methane variations in the past two decades using measurements of CH₄ mixing ratio and isotopic composition. *Atmos. Chem. Phys.* **11**, 9141–9153 (2011).
- Bousquet, P. *et al.* Contribution of anthropogenic and natural sources to atmospheric methane variability. *Nature* **443**, 439–443 (2006).
- Neef, L., van Weele, M. & van Velthoven, P. Optimal estimation of the present-day global methane budget. *Glob. Biogeochem. Cycles* **24**, GB4024 (2010).
- Wahlen, M., Tanaka, N., Henry, R. & Yoshinari, T. ¹³C, D and ¹⁴C in methane. *Eos* **68**, 1220 (1987).
- Fisher, R. E. *et al.* Arctic methane sources: Isotopic evidence for atmospheric inputs. *Geophys. Res. Lett.* **38**, L21803 (2011).
- Kepler, F., Hamilton, J. T. G., Brass, M. & Rockmann, T. Methane emissions from terrestrial plants under aerobic conditions. *Nature* **439**, 187–191 (2006).
- Nisbet, R. E. R. *et al.* Emission of methane from plants. *Proc. R. Soc. B-Biol. Sci.* **276**, 1347–1354 (2009).
- Curry, C. L. Modeling the soil consumption of atmospheric methane at the global scale. *Glob. Biogeochem. Cycles* **21**, GB4012 (2007).
- Zhuang, Q. *et al.* Methane fluxes between terrestrial ecosystems and the atmosphere at northern high latitudes during the past century: A retrospective analysis with a process-based biogeochemistry model. *Glob. Biogeochem. Cycles* **18**, GB3010 (2004).
- Allan, W., Struthers, H. & Lowe, D. C. Methane carbon isotope effects caused by atomic chlorine in the marine boundary layer: Global model results compared with Southern Hemisphere measurements. *J. Geophys. Res.: Atmos.* **112**, D04306 (2007).
- Naik, V. *et al.* Preindustrial to present day changes in tropospheric hydroxyl radical and methane lifetime from the Atmospheric Chemistry and Climate Model Intercomparison Project (ACCMIP). *Atmos. Chem. Phys.* **13**, 5277–5298 (2013).
- Bastviken, D., Tranvik, L. J., Downing, J. A., Crill, P. M. & Enrich-Prast, A. Freshwater methane emissions offset the continental carbon sink. *Science* **331**, 50 (2011).
- Walter, K. M., Smith, L. C. & Stuart Chapin, F. Methane bubbling from northern lakes: present and future contributions to the global methane budget. *Phil. Trans. R. Soc. A* **365**, 1657–1676 (2007).
- Huang, J. & Prinn, R. G. Critical evaluation of emissions of potential new gases for OH estimation. *J. Geophys. Res.* **107**, 4784 (2002).
- Montzka, S. A. *et al.* Small interannual variability of global atmospheric hydroxyl. *Science* **331**, 67–69 (2011).
- Etioppe, G., Lassey, K. R., Klusman, R. W. & Boschi, E. Reappraisal of the fossil methane budget and related emission from geologic sources. *Geophys. Res. Lett.* **35**, L09307 (2008).
- Shakhova, N. *et al.* Extensive methane venting to the atmosphere from sediments of the East Siberian Arctic Shelf. *Science* **327**, 1246 (2010).
- Lassey, K. R., Lowe, D. C. & Smith, A. M. The atmospheric cycling of radiomethane and the “fossil fraction” of the methane source. *Atmos. Chem. Phys.* **7**, 2141–2149 (2007).
- Voulgarakis, A. *et al.* Analysis of present day and future OH and methane lifetime in the ACCMIP simulations. *Atmos. Chem. Phys.* **13**, 2563–2587 (2013).
- Melton, J. R. *et al.* Present state of global wetland extent and wetland methane modelling: conclusions from a model intercomparison project (WETCHIMP). *Biogeosciences* **10**, 753–788 (2013).
- Hodson, E. L., Poulter, B., Zimmermann, N. E., Prigent, C. & Kaplan, J. O. The El Niño Southern Oscillation and wetland methane interannual variability. *Geophys. Res. Lett.* **38**, L08810 (2011).
- Ringeval, B. *et al.* Climate-CH₄ feedback from wetlands and its interaction with the climate-CO₂ feedback. *Biogeosciences* **8**, 2137–2157 (2011).
- Spahni, R. *et al.* Constraining global methane emissions and uptake by ecosystems. *Biogeosciences* **8**, 1643–1665 (2011).
- Riley, W. J. *et al.* Barriers to predicting changes in global terrestrial methane fluxes: analyses using CLM4Me, a methane biogeochemistry model integrated in CESM. *Biogeosciences* **8**, 1925–1953 (2011).
- Mastrandrea, M. D. *et al.* *Guidance Note for Lead Authors of the IPCC Fifth Assessment Report on Consistent Treatment of Uncertainties* (IPCC, 2010); <http://www.ipcc.ch>.
- Ohara, T. *et al.* An Asian emission inventory of anthropogenic emission sources for the period 1980–2020. *Atmos. Chem. Phys.* **7**, 4419–4444 (2007).
- Bergamaschi, P. *et al.* Inverse modeling of global and regional CH₄ emissions using SCIAMACHY satellite retrievals. *J. Geophys. Res.* <http://dx.doi.org/10.1029/2009JD012287> (2009).
- Van der Werf, G. R. *et al.* Global fire emissions and the contribution of deforestation, savanna, forest, agricultural, and peat fires (1997–2009). *Atmos. Chem. Phys.* **10**, 11707–11735 (2010).
- Sanderson, M. G. Biomass of termites and their emissions of methane and carbon dioxide: A global database. *Glob. Biogeochem. Cycles* **10**, 543–557 (1996).
- Bousquet, P., Hauglustaine, D. A., Peylin, P., Carouge, C. & Ciais, P. Two decades of OH variability as inferred by an inversion of atmospheric transport and chemistry of methyl chloroform. *Atmos. Chem. Phys.* **5**, 2635–2656 (2005).

50. Simpson, I. J., Rowland, F. S., Meinardi, S. & Blake, D. R. Influence of biomass burning during recent fluctuations in the slow growth of global tropospheric methane. *Geophys. Res. Lett.* **33**, L22808 (2006).
51. Dlugokencky, E. J. *et al.* Changes in CH₄ and CO growth rates after the eruption of Mt Pinatubo and their link with changes in tropical tropospheric UV flux. *Geophys. Res. Lett.* **23**, 2761–2764 (1996).
52. Langenfelds, R. L. *et al.* Interannual growth rate variations of atmospheric CO₂ and its delta ¹³C, H₂, CH₄, and CO between 1992 and 1999 linked to biomass burning. *Glob. Biogeochem. Cycles* **16**, 1048 (2002).
53. Bousquet, P. *et al.* Source attribution of the changes in atmospheric methane for 2006–2008 *Atmos. Chem. Phys.* **11**, 3689–3700 (2011).
54. Dlugokencky, E. J., Masarie, K. A., Lang, P. M. & Tans, P. P. Continuing decline in the growth rate of the atmospheric methane burden. *Nature* **393**, 447–450 (1998).
55. Environmental Protection Agency. Global Anthropogenic Non-CO₂ Greenhouse Gas Emissions: 1990–2030. (US Environmental Protection Agency, 2011).
56. European Commission, Joint Research Centre/Netherlands Environmental Assessment Agency. Emission Database for Global Atmospheric Research (EDGAR) (version 4.2) (2011); <http://edgar.jrc.ec.europa.eu>.
57. Aydin, M. *et al.* Recent decreases in fossil-fuel emissions of ethane and methane derived from firm air. *Nature* **476**, 198–201 (2011).
58. Kai, F. M., Tyler, S. C., Randerson, J. T. & Blake, D. R. Reduced methane growth rate explained by decreased Northern Hemisphere microbial sources. *Nature* **476**, 194–197 (2011).
59. Levin, I. *et al.* No inter-hemispheric δ¹³CH₄ trend observed. *Nature* **486**, E3–E4 (2012).
60. Bloom, A. A., Palmer, P. I., Fraser, A., Reay, D. S. & Frankenberg, C. Large-scale controls of methanogenesis inferred from methane and gravity spaceborne data. *Science* **327**, 322–325 (2010).
61. Prigent, C., Papa, F., Aires, F., Rossow, W. B. & Matthews, E. Global inundation dynamics inferred from multiple satellite observations, 1993–2000. *J. Geophys. Res.* **112**, D12107 (2007).
62. FLUXNET database; <http://fluxnet.ornl.gov>.
63. Lewis, S. L., Brando, P. M., Phillips, O. L., van der Heijden, G. M. F. & Nepstad, D. The 2010 Amazon drought. *Science* **331**, 554–554 (2011).
64. Houweling, S. *et al.* Iconic CO₂ Time Series at Risk. *Science* **337**, 1038–1040 (2012).
65. Lamarque, J. F. *et al.* The Atmospheric Chemistry and Climate Model Intercomparison Project (ACCMIP): overview and description of models, simulations and climate diagnostics. *Geosci. Model Dev.* **6**, 179–206 (2013).
66. Patra, P. K. *et al.* TransCom model simulations of CH₄ and related species: linking transport, surface flux and chemical loss with CH₄ variability in the troposphere and lower stratosphere. *Atmos. Chem. Phys.* **11**, 12813–12837 (2011).
67. Kiemle, C. *et al.* Sensitivity studies for a space-based methane lidar mission. *Atmos. Meas. Tech.* **4**, 2195–2211 (2011).
68. Shindell, D. *et al.* Simultaneously mitigating near-term climate change and improving human health and food security. *Science* **335**, 183–189 (2012).
69. Howarth, R., Santoro, R. & Ingraffea, A. Methane and the greenhouse-gas footprint of natural gas from shale formations. *Climatic Change* **106**, 679–690 (2011).
70. Cathles, L., Brown, L., Taam, M. & Hunter, A. A commentary on “The greenhouse-gas footprint of natural gas in shale formations” by R. W. Howarth, R. Santoro, and Anthony Ingraffea. *Climatic Change* **113**, 525–535 (2012).
71. Koven, C. D. *et al.* Permafrost carbon-climate feedbacks accelerate global warming. *Proc. Natl Acad. Sci. USA* **108**, 14769–14774 (2011).
72. Global Carbon Project (2013); <http://www.globalcarbonproject.org/index.htm>.
73. Bruhwiler, L., Dlugokencky, E. J. & Masarie, K. *AGU Fall Meeting* abstr. B11G-01 (2011).
74. Chen, Y. H. & Prinn, R. G. Estimation of atmospheric methane emissions between 1996 and 2001 using a three-dimensional global chemical transport model. *J. Geophys. Res.* **111**, D10307 (2006).
75. Fraser, A. *et al.* Estimating regional methane surface fluxes: the relative importance of surface and GOSAT mole fraction measurements. *Atmos. Chem. Phys.* **13**, 5697–5713 (2013).
76. Hein, R., Crutzen, P. J. & Heimann, M. An inverse modeling approach to investigate the global atmospheric methane cycle. *Glob. Biogeochem. Cycles* **11**, 43–76 (1997).
77. Pison, I., Bousquet, P., Chevallier, F., Szopa, S. & Hauglustaine, D. Multi-species inversion of CH₄, CO and H₂ emissions from surface measurements. *Atmos. Chem. Phys.* **9**, 5281–5297 (2009).
78. Mieville, A. *et al.* Emissions of gases and particles from biomass burning during the 20th century using satellite data and an historical reconstruction. *Atmos. Environ.* **44**, 1469–1477 (2010).
79. Van het Bolscher, M. *et al.* Emission data sets and methodologies for estimating emissions (eds Schultz, M. G. & Rast, S.) (2007).
80. Wiedinmyer, C. *et al.* The Fire INventory from NCAR (FINN): a high resolution global model to estimate the emissions from open burning. *Geosci. Model Dev.* **4**, 625–641 (2011).
81. Dentener, F. *et al.* The impact of air pollutant and methane emission controls on tropospheric ozone and radiative forcing: CTM calculations for the period 1990–2030 *Atmos. Chem. Phys.* **5**, 1731–1755 (2005).
82. Environmental Protection Agency. Methane and Nitrogen Oxide Emissions From Natural Sources. (US Environmental Protection Agency, 2010).
83. Williams, J. E., Strunk, A., Huijnen, V. & van Weele, M. The application of the Modified Band Approach for the calculation of on-line photodissociation rate constants in TM5: implications for oxidative capacity. *Geosci. Model Dev.* **5**, 15–35 (2012).
84. Gurney, K. R. *et al.* Transcom 3 inversion intercomparison: Model mean results for the estimation of seasonal carbon sources and sinks. *Glob. Biogeochem. Cycles* **18**, GB2010 (2004).
85. Prinn, R. G. *et al.* Evidence for substantial variations of atmospheric hydroxyl radicals in the past two decades. *Science* **292**, 1882–1888 (2001).
86. Beck, V. *et al.* Methane airborne measurements and comparison to global models during BARCA. *Journal of Geophysical Research: Atmospheres* **117**, D15310 (2012).
87. Dickens, G. R. Methane hydrates in quaternary climate change — The clathrate gun hypothesis. *Science* **299**, 1017–1017 (2003).
88. Hoelzemann, J. J., Schultz, M. G., Brasseur, G. P., Granier, C. & Simon, M. Global Wildland Fire Emission Model (GWEM): Evaluating the use of global area burnt satellite data. *J. Geophys. Res.* **109**, D14S04 (2004).
89. Ito, A. & Penner, J. E. Global estimates of biomass burning emissions based on satellite imagery for the year 2000. *J. Geophys. Res.* **109**, D14S05 (2004).
90. Rhee, T. S., Kettle, A. J. & Andreae, M. O. Methane and nitrous oxide emissions from the ocean: A reassessment using basin-wise observations in the Atlantic. *J. Geophys. Res.* **114**, D12304 (2009).
91. Sugimoto, A., Inoue, T., Kirtibutr, N. & Abe, T. Methane oxidation by termite mounds estimated by the carbon isotopic composition of methane. *Glob. Biogeochem. Cycles* **12**, 595–605 (1998).
92. Kasibhatla, K. *et al.* *Inverse Methods in Global Biogeochemical Cycles*, Volume 114 (AGU, 2000).
93. Rodgers, C. D. *Inverse Methods for Atmospheric Sounding: Theory and Practice* (World Scientific, 2000).
94. Krol, M. & Lelieveld, J. Can the variability in tropospheric OH be deduced from measurements of 1,1,1-trichloroethane (methyl chloroform)? *J. Geophys. Res.* **108**, 4125 (2003).

Acknowledgements

This paper is the result of an international collaboration of scientists organized by the Global Carbon Project, a joint project of the Earth System Science Partnership. This work was supported by: the UK NERC National Centre for Earth Observation; the European Commission's 7th Framework Programme (FP7/2007-2013) projects MACC (grant agreement no. 218793) and GEOCARBON (grant agreement no. 283080); contract DE-AC52-07NA27344 with different parts supported by the US DOE IMPACTS and SciDAC Climate Consortium projects; computing resources of NERSC, which is supported by the US DOE under contract DE-AC02-05CH11231; NOAA flask data for CH₂Cl₂ (made available by S. Montzka); the Australian Climate Change Science Program, and ERC grant 247349. Simulations from LSCE were performed using HPC resources from DSM-CCRT and CCRT/CINES/IDRIS under the allocation 2012-t2012012201 made by GENCI (Grand Equipement National de Calcul Intensif). We thank the EDGAR group at JRC (Italy) and US-EPA for providing estimates of anthropogenic emissions.

Author contributions

S.K., P. Bousquet, P.C., J.G.C. and C.L.Q. designed the study and provided conceptual advice. S.K., P. Bousquet and M. Saunio processed data sets, developed figures and wrote the manuscript. E.J.D., M. Schmidt, P.J.F., P.B.K., L.P.S., R.L.L., R.G.P., M.R., R.F.W., D.R.B. and I.J.S. provided atmospheric *in situ* data. P. Bousquet, P. Bergamaschi, L.B., F.C., L.F., A.F., S.H., P.I.P. and I.P. provided top-down inversion results (all five emission categories). S.C., E.L.H., B.P., B.R., M. Santini, R.S. and G.R.v.d.W. provided bottom-up modelling and inventory data sets for wetland, biomass burning and termite emissions. D.B., P.C.-S., B.J., J.-F.L., V.N., D.P., D.T.S., S.A.S., K.S., S.S., A.V., M.v.W., J.E.W. and G.Z. provided bottom-up estimates of CH₄ loss due to OH. All authors contributed extensively to the work presented in this paper, and to revisions of the manuscript.

Additional information

Reprints and permissions information is available online at www.nature.com/reprints. Correspondence should be addressed to philippe.bousquet@lscce.ipsl.fr.

Competing financial interests

The authors declare no competing financial interests.

Stefanie Kirschke¹, Philippe Bousquet^{1*}, Philippe Ciais¹, Marielle Saunois¹, Josep G. Canadell², Edward J. Dlugokencky³, Peter Bergamaschi⁴, Daniel Bergmann⁵, Donald R. Blake⁶, Lori Bruhwiler³, Philip Cameron-Smith⁵, Simona Castaldi^{7,8}, Frédéric Chevallier¹, Liang Feng⁹, Annemarie Fraser⁹, Martin Heimann¹⁰, Elke L. Hodson¹¹, Sander Houweling^{12,13}, Béatrice Josse¹⁴, Paul J. Fraser¹⁵, Paul B. Krummel¹⁵, Jean-François Lamarque¹⁶, Ray L. Langenfelds¹⁵, Corinne Le Quéré¹⁷, Vaishali Naik¹⁸, Simon O'Doherty¹⁹, Paul I. Palmer⁹, Isabelle Pison¹, David Plummer²⁰, Benjamin Poulter¹, Ronald G. Prinn²¹, Matt Rigby²², Bruno Ringeval^{13,23,24}, Monia Santini⁸, Martina Schmidt¹, Drew T. Shindell²⁵, Isobel J. Simpson⁶, Renato Spahni²⁶, L. Paul Steele¹⁵, Sarah A. Strode^{27,28}, Kengo Sudo²⁹, Sophie Szopa¹, Guido R. van der Werf³⁰, Apostolos Voulgarakis^{25,31}, Michiel van Weele³², Ray F. Weiss³³, Jason E. Williams³² and Guang Zeng³⁴

¹LSCE-CEA-UVSQ-CNRS, Orme des Merisiers, 91190 Gif-sur-Yvette, France, ²Global Carbon Project, CSIRO Marine and Atmospheric Research, GPO Box 3023, Canberra, ACT 2601, Australia, ³NOAA ESRL, 325 Broadway, Boulder, Colorado 80305, USA, ⁴Institute for Environment and Sustainability, Joint Research Centre, TP290, I-21027 Ispra (Va), Italy, ⁵Lawrence Livermore National Laboratory, PO Box 808, Livermore, California 94551-0808, USA, ⁶University of California Irvine, 570 Rowland Hall, Irvine, California 92697, USA, ⁷Department of Environmental Sciences, Second University of Naples, via Vivaldi 43, 81100 Caserta, Italy, ⁸Centro Euro-Mediterraneo sui Cambiamenti Climatici (CMCC), via Augusto Imperatore 16, 73100 Lecce, Italy, ⁹School of Geosciences, University of Edinburgh, Crew Building, West Mains Road, Edinburgh, EH9 3JN, UK, ¹⁰Max Planck Institute for Biogeochemistry, PF 100164, D-07701 Jena, Germany, ¹¹Swiss Federal Research Institute WSL, Zuercherstrasse 111, 8903 Mensdorf, Switzerland, ¹²SRON, Netherlands Institute for Space Research, Sorbonnelaan 2, 3584 CA Utrecht, The Netherlands, ¹³Institute for Marine and Atmospheric Research Utrecht, Sorbonnelaan 2, 3584 CA Utrecht, The Netherlands, ¹⁴Météo France, CNRM/GMGEC/CARMA, 42 av. G. Coriolis, 31057 Toulouse, France, ¹⁵Centre for Australian Weather and Climate Research/CSIRO Marine and Atmospheric Research, Aspendale, Victoria 3195, Australia, ¹⁶NCAR, PO Box 3000, Boulder, Colorado 80307-3000, USA, ¹⁷Tyndall Centre for Climate Change Research, University of East Anglia, Norwich Research Park, Norwich, NR4 7TJ, UK, ¹⁸UCAR/NOAA Geophysical Fluid Dynamics Laboratory, 201 Forrestal Road, Princeton, New Jersey 08540, USA, ¹⁹University of Bristol, Office Old Park Hill, Cantock's Close, Clifton, Bristol BS8 1TS, UK, ²⁰Canadian Centre for Climate Modelling and Analysis, Environment Canada, 550 Sherbrooke Street West, West Tower, Montréal, Quebec, H3A 1B9, Canada, ²¹Massachusetts Institute of Technology, Building 54-1312, Cambridge, Massachusetts 02139-2307, USA, ²²School of Chemistry, University of Bristol, Cantocks Close, Bristol BS8 1TS, UK, ²³IMAU, Utrecht University, Princetonplein 5, 3584 CC Utrecht, The Netherlands, ²⁴Department of Systems Ecology, VU University Amsterdam, De Boelelaan 1085, 1081 HV Amsterdam, The Netherlands, ²⁵NASA Goddard Institute for Space Studies, 2880 Broadway, New York, New York 10025, USA, ²⁶University of Bern, Physics Institute, Climate and Environmental Physics, Sidlerstrasse 5, CH-3012 Bern, Switzerland, ²⁷NASA Goddard Space Flight Centre, Greenbelt, Maryland 20771, USA, ²⁸Universities Space Research Association, NASA Goddard Space Flight Centre, Greenbelt, Maryland 20771, USA, ²⁹Graduate School of Environmental Studies, Nagoya University Furo-cho, Chikusa-ku, Nagoya 464-8601, Japan, ³⁰VU University, Faculty of Earth and Life Sciences, De Boelelaan 1085, 1081 HV Amsterdam, The Netherlands, ³¹Department of Physics, Imperial College London, London SW7 2AZ, UK, ³²Royal Netherlands Meteorological Institute (KNMI), PO Box 201, 3730 AE De Bilt, The Netherlands, ³³Scripps Institution of Oceanography, UCSD, La Jolla, California 92093-0244, USA, ³⁴National Institute of Water and Atmospheric Research, Private Bag 50061, Omakau, Central Otago 9352, New Zealand. *e-mail: Philippe.Bousquet@lsce.ipsl.fr